

BULLETIN

Corpus Christi Geological Society



and

Coastal Bend Geophysical Society



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CORPUS CHRISTI GEOLOGICAL SOCIETY
P.O. BOX 1068* C.C.TX. 78403
2019-2020

www.ccgeo.org

OFFICERS

President	Austin Nye	361-452-1435	austin@nyexp.us
President Elect			
Vice President	Randy Bissell	361-885-0110	randyb@headingtonenergy.com
Secretary	Emily Olson	361-542-7434	emily@olsongeo.com
Treasurer	Sebastian Wiedmann	361-946-4430	swiedmann.geo@gmail.com
Past President	Frank Cornish	361-563-9184	frank.cornish@gmail.com
Councilor I	Rick Paige	361-884-8824	rpaigesio2@gmail.com
Councilor II	BJ Thompson	361-816-2326	william.thompson426@gmail.com

AAPG DELEGATE

EDITORS

Bulletin Editor	Marian Wiedmann	361-855-2542	wiedgulf@gmail.com
Bulletin Tech. Editor	Jeanie Timmermann	361-883-1492	jeanie@industrialcorrosion.net
Web Master	Josh Pollard	361-654-3100	support@interconnect.net

GEOLOGICAL SOCIETY COMMITTEES & CHAIRPERSONS

MEMBERS PLEASE NOTE COMMITTEES CONTINUE TO FUNCTION WITHOUT A NAMED CHAIR, ALL ARE WELCOME TO ANY COMMITTEE THAT SUITES YOUR INTEREST

Advertising/Bus. Cards Arrangements	Robby Sterett Wes Gisler & Will Graham	361-739-5618 830-239-4651 361-885-0110	robert.sterett@gmail.com arrangements@ccgeo.org willg@headingtonenergy.com
Continuing Education			
Education	Dawn Bissell	361-960-2151	bissells@swbell.net
Scholarship Program	Dawn Bissell	361-960-2151	bissells@swbell.net
Fishing Tournament	Leighton Devine	361-882-8400	ldevine@suemaur.com
History	OPEN (Ray Govett)	361-855-0134	ray30@hotmail.com
Membership	Dorothy Jordan Randy Bissell	361-885-0110 361-885-0113	dorothyj@headington.com randyb@headington.com
Type Logs			



COASTAL BEND GEOPHYSICAL SOCIETY
P.O. BOX 2741*C.C. TX. 78403
2017-2018

OFFICERS

President	Dr. Subbarao Yelisetti	361-593-4894	subbarao.yelisetti@tamuk.edu
Vice President	Samara Omar	361-902-2826-office 361-446-7334-mobile	Samara_Omar@eogresources.com
Secretary/Treasurer	Erik Scott	361-902-2880-office 361-800-3364-mobile	Erik_Scott@eogresources.com
Golf Chairman	Mark Wiley	361-902-2844-office 361-445-6712-mobile	Mark_Wiley@eogresources.com
Scholarship Chairman	Matt Hammer	361-888-4792 361-563-6137	mhammer@royalcctx.com

**Visit the geological
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CCGS/CBGS JOINT MEETING SCHEDULE 2019-2020

September 2019							October 2019							November 2019							
S	M	T	W	Th	F	S	S	M	T	W	Th	F	S	S	M	T	W	Th	F	S	
1	2	3	4	5	6	7			1	2	3	4	5							1	2
8	9	10	11	12	13	14	6	7	8	9	10	11	12	3	4	5	6	7	8	9	
15	16	17	18	19	20	21	13	14	15	16	17	18	19	10	11	12	13	14	15	16	
22	23	24	25	26	27	28	20	21	22	23	24	25	26	17	18	19	20	21	22	23	
29	30						27	28	29	30	31			24	25	26	27	28	29	30	

Thursday, Sept. 26th at 5:30-8:00p.m. Kickoff at Hoegemeyer's Barbeque Barn.

11:30-1:00 pm
Speaker: Dr. Neil Bockoven

December 2019							January 2020							February 2020						
S	M	T	W	Th	F	S	S	M	T	W	Th	F	S	S	M	T	W	Th	F	S
1	2	3	4	5	6	7				1	2	3	4							1
8	9	10	11	12	13	14	5	6	7	8	9	10	11	2	3	4	5	6	7	8
15	16	17	18	19	20	21	12	13	14	15	16	17	18	9	10	11	12	13	14	15
22	23	24	25	26	27	28	19	20	21	22	23	24	25	16	17	18	19	20	21	22
29	30	31					26	27	28	29	30	31		23	24	25	26	27	28	29

CCGS/CBGS Joint Meeting Schedule 2019-2020

March 2020							April 2020							May 2020						
S	M	T	W	Th	F	S	S	M	T	W	Th	F	S	S	M	T	W	Th	F	S
1	2	3	4	5	6	7				1	2	3	4						1	2
8	9	10	11	12	13	14	5	6	7	8	9	10	11	3	4	5	6	7	8	9
15	16	17	18	19	20	21	12	13	14	15	16	17	18	10	11	12	13	14	15	16
22	23	24	25	26	27	28	19	20	21	22	23	24	25	17	18	19	20	21	22	23
29	30	31					26	27	28	29	30			24	25	26	27	28	29	30
														31						

Calendar of Meetings and Events

Calendar of Area Monthly Meetings

Corpus Christi Geological/Geophysical Society.....	Third Wed.—11:30a.m.
SIPES Corpus Christi Luncheons.....	Last Tues.—11:30a.m.
South Texas Geological Society Luncheons.....	Second Wed—noon San Antonio
San Antonio Geophysical Society Meetings.....	Fourth Tuesday
Austin Geological Society.....	First Monday
Houston Geological Society Luncheons.....	Last Wednesday
Central Texas Section of Society of Mining, Metallurgy & Exp.....	2 nd Tues every other month in San Antonio



CCGS PRESIDENT'S LETTER

Welcome to the 2019/2020 season of our illustrious society. We have some changes this year to the luncheon schedule; only four typical luncheons, and two technical sessions courtesy of the CBGS. Our BBQ kickoff will be held at Hoegemeyer's Barbeque Barn on Thursday September 26, 2019 at 5:30pm to 8:00pm. Our regular luncheons will be held in November 20th, January 15th, April 15th and May 20th. The two technical sessions will be held in October and February.

Our fishing tournament this year was a big success, raising over \$2000.00 for the scholarship fund. Over the years we have donated \$55,000 to the CCGS scholarship fund.

Thank you to our current board, Randy Bissel, Vice President; Sebastian Wiedmann, treasurer; Emily Olson, Secretary; BJ Thompson, Counselor; Rick Paige, Counselor. As always, we need more volunteers. A special thank you to Emily Olson for stepping up, and Brianna Watkins for offering to manage our social media foot print.

Looking forward to a great year.

Austin Nye—VP Exploration & Land
CCGS President



CBGS President's Letter

CBGS Board 2019-2020

Dr. Subbarao Yelisetti- President

Samara Omar- Vice President

Erik Scott- Secretary/ Treasurer

Matt Hammer - Scholarship Chair

Mark Wiley - Golf Chair

Education – Robert Schneider

CBGS Scholarships

The board awarded three scholarships of \$2,000 each to undergraduate geophysics majors from Texas A&M University-College Station, University of Houston and Texas A&M University-Kingsville in 2017-2018. We will be awarding the scholarships again this year.

Scholarship Requirements

Criteria for awarding the Scholarship from Coastal Bend Geophysical Society of Corpus Christi, Texas:

1. Scholarships are open to undergraduate or graduate students.
2. Must have declared major in Geophysics, or Geology with a concentration in Geophysics or Petrophysics.
3. Preference is given to students attending Coastal Bend schools (TAMU-K, TAMU-CC and Del Mar College), then to Coastal Bend natives attending other universities.
4. Must have a GPA of at least 3.0 and be in good standing with the school.
5. Must make effort to attend a Coastal Bend Geophysical Society Meeting in Corpus Christi Texas after being awarded a scholarship to be recognized by the society.

News

- According to the U.S. Energy Information Administration (EIA) projections, the U.S. oil production would rise to 12.27 million bpd in 2019 from a record 10.99 million bpd in 2018.
- According to EIA projections, U.S. crude output would rise by 85,000 barrels per day (bpd) in September to a record 8.77 million bpd.
- At the time of writing this report, U.S. crude futures were trading around \$54 a barrel for the balance of 2019 and \$52 in calendar 2020 as reported by Scott DiSavino on reuters.com.

- The U.S. active rig count so far this year has averaged 1,004.
- According to the analysts at Simmons & Co forecast, the average combined oil and gas rig count will slide from a four-year high of 1,032 in 2018 to 970 in 2019 and 955 in 2020 before rising to 997 in 2021 as reported by Scott DiSavino on reuters.com.

CBGS Business

CBGS currently has 52 active members, 4 honorary members, and 60 student members.

CBGS workshops/talks

- As part of the annual Kickoff Bar-B-Q, CBGS hosted Dr. Satinder Chopra on Sep 19, 2018 at the EOG conference center. His talk was entitled “Seismic reservoir characterization of Utica-Point Pleasant shale with efforts at quantitative interpretation – a case study”.
- CBGS offered a land seismic acquisition workshop on Dec 5th in EOG conference center with the following talks.

Talks:

1. A Brief Introduction to Seismic Acquisition (Students & New Professionals Encouraged) by Lonnie Blake, EOG
2. A Comparison of Long, Short, and Slip Sweep 3D Data Image Volumes Acquired And Constrained By Equivalent Source Time (KWP Phase I) by J. W. (Tom) Thomas, Kevin Werth, Tom Phillips, Chris Lindsey Dawson Geophysical Co.
3. A Comparison of 3D Multi-Component (9C) Data Image Volumes Acquired With Conventional and Simultaneous Source Techniques Also With Adequate Spatial Resolution For Compressive Sensing Investigation (KWP Phase 2) by J. W. (Tom) Thomas, Kevin Werth, Tom Phillips, Chris Lindsey Dawson Geophysical Co.
4. SAExporation - More, Recent Advances in Onshore Seismic Data Acquisition Methods by Howard Watt, SAExporation, Houston

CBGS is looking forward to offer many such workshops in the future. Topic/speaker suggestions are welcome. Email your suggestions to Samara_Omar@eogresources.com or Subbarao.Yelisetti@tamuk.edu

Golf Tournament

CBGS organized its annual **Golf Tournament** to fund its scholarship program in the first week of October, 2018 at Northshore Country Club. Raised ~\$4,000 for the scholarship fund.

If you are interested in our next Golf Tournament, please contact Mark Wiley at Mark_Wiley@eogresources.com

New Degree Tracks at TAMUK

- Texas A&M University-Kingsville (TAMUK) started its first cohort of MS Petrophysics program in Fall 2018. If you are interested in joining this program in Fall 2019, please contact the graduate coordinator for MS in Petrophysics, Dr. Subbarao Yelisetti at Subbarao.Yelisetti@tamuk.edu.
- **BS degree in Geophysics, Minor in Geophysics and Certification in Geophysics** offered at Texas A&M University-Kingsville since Fall 2017. Interested students can contact Dr. Subbarao Yelisetti (Subbarao.Yelisetti@tamuk.edu) for additional information.

Exploration Geophysics and Borehole Geophysics classes

PHYS 5382 Exploration Geophysics and **PHYS 5388 Borehole Geophysics** classes are offered in Fall 2019 at Texas A&M University-Kingsville. This is available for the professional community as well as our students. You can sign up as a “transient” student in order to take classes without actually enrolling in the school. If anyone in the professional community wishes to sign up for this, please contact, Dr. Subbarao Yelisetti Subbarao.Yelisetti@tamuk.edu.

SEG Distinguished Lecture

CBGS and TAMUK SEG student chapter organized 2018 SEG Distinguished Lecture in January, 2018. We wish to organize many more lectures in the future.

Education/Events

-SEG

SEG 2019 annual meeting will be held in San Antonio, TX from Sep 15-20th. Abstract deadline is April 1, 2019. See <https://seg.org/Annual-Meeting-2019> for additional details.

See <https://seg.org/Education/Lectures/Distinguished-Lectures> for information about upcoming SEG distinguished lecture in Houston and other locations.

See <https://seg.org/Education/Lectures/Honorary-Lectures> for SEG honorary lecture locations in Texas.

-AGU

2019 Fall AGU annual meeting will be held in San Francisco, CA from December 9-13th, 2019. <https://fallmeeting.agu.org/2018/future-meetings/>

-GSA

The Geological Society of America’s 131st annual meeting will be held in Phoenix, AZ from 22-25th September 2019.

http://www.geosociety.org/GSA/Events/Annual_Meeting/GSA/Events/gsa2019.aspx

Monthly Saving

"Striking oil would seem to be rather an easy process on such grounds. It would be worth trying"
- Malcolm McLeod (1872) commenting on the journal of Chief Factor Archibald McDonald
written in 1828 in reference to the Bituminous Springs along the Athabasca River.

Monthly Summary

Texas Oil and Gas Info	Current Month	Last Month	Difference	
Texas Production	MMBO/BCF	MMBO/BCF	MMBO/BCF	
Oil	125.7	134.6	-8.9	April
Condensate	15.4	15.3	0.1	April
Gas	743.9	773.4	-29.5	April
	Current Month	Yr to date - 2019	Yr to date - 2018	
Texas Drilling Permits	912	7166	8330	July
Oil wells	212	1746	2160	July
Gas wells	58	436	519	July
Oil and Gas wells	537	4416	5075	July
Other	9	69	80	July
Total Completions	699	5749	6514	July
Oil Completions	499	4278	5089	July
Gas Completions	156	1181	1084	July
New Field Discoveries	0	7	15	July
Other	2	18	24	July

Subbarao Yelisetti
President, CBGS



The Corpus Christi Geological Society
The Coastal Bend Geophysical Society



Invite you to the
Society Kickoff Bar-B-Q

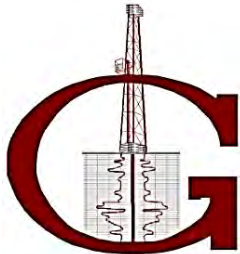
Thursday, September 26, 2019
5:30 to 8:00 p.m.

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RSVP (required) by Monday, Sept. 23

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Wes Gisler, wes@gislerbrotherslogging.com, 1-830-239-4651

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Memorial
Barnard Dietz



Upon hearing that Barnard Dietz had passed, I found myself thinking back as to when I first became acquainted with him. In 1966 I came to work for Texas Oil & Gas as an engineer in Corpus Christi. Barnard was the Gulf Coast District Geologist for TXO. The next year he was promoted to Chief Geologist, assigned to the corporate office in Dallas. All major geological prospects had to receive his blessing prior to proceeding further in the planning stage. Presenting a Gulf Coast prospect to Barnard was sometimes an uncomfortable job. He already had strong opinions about the geology and was very outspoken, and rightfully so. One main thing I admired about him was even though some discussions could become very heated at times, he never seemed to hold a grudge. Another thing I admired was that every time we spoke, he would always offer me a word of encouragement. He was definitely a team player.

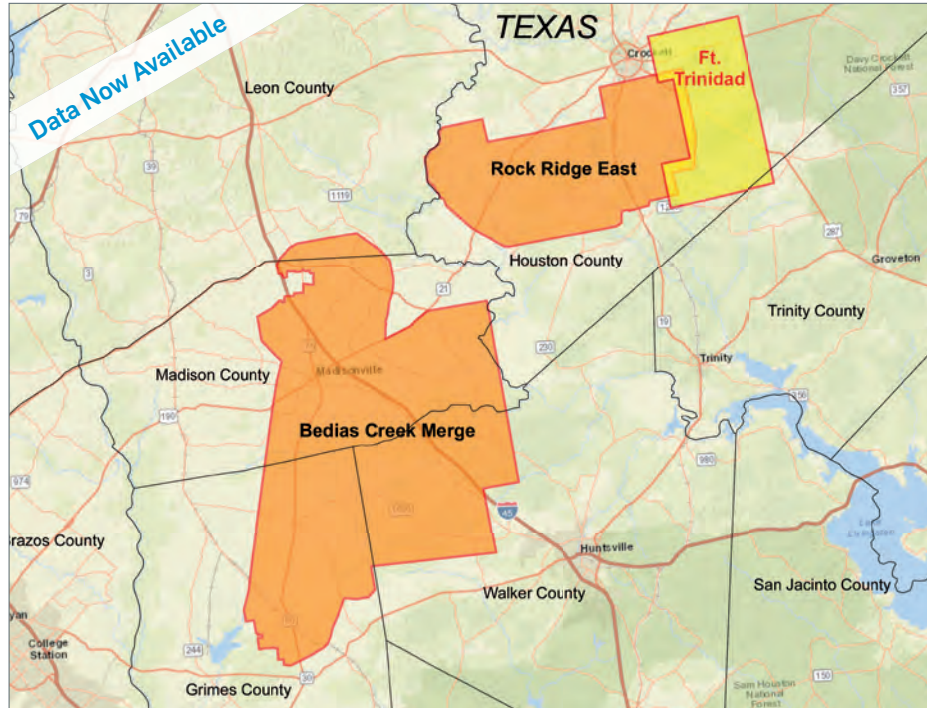
One major parameter defining the success of an individual, is whether or not they lived their life in such a way as to make a positive impact on those they had contact with. Barnard Dietz did that. I'm very honored to have been a friend and a colleague of his. May he rest in peace.

Louis Little



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Scott Tinley
+1 832 351 8544
scott.tinley@cgg.com

Cheryl Oxsheer
+1 832 351 8463
cheryl.oxsheer@cgg.com



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Annual CCGS Fishing Tournament



The 10th Annual CCGS Fishing Tournament was a huge success. The festivities started Thursday evening with the Dock Party/Captain's Meeting at Treasure Island Bar in beautiful Port Aransas, Texas. Saltwater Grill provided some great food followed by the Captain's Meeting and door prize giveaway. Thank you to all the volunteers who made our Thursday event run so smoothly.

Eighteen inshore teams set out Friday morning under perfect conditions, and brought back some really impressive fish. We had a very competitive group of anglers making the afternoon weigh-in a great experience. Thanks to Pete Graham and Schlumberger for the fajitas on Friday.

The last 10 years have been a great experience for all of us involved in putting this tournament together. With the help of our tremendous sponsors, most of which have been with us from day one, our winners took home over \$15,000.00 in cash and prizes. To date we have donated over \$55,000.00 to support the CCGS Scholarship Trust Fund and various other continuing education funds. I look forward to the coming years and hope we can continue to put on a successful industry event.

Leighton Devine
Fishing Tournament Chairman



Kelby Broussard (TEXEGY)



**David Musgrove
(TEXLA Energy Management)**



Zach Corcoran (Sir Drake Oil & Gas Company)

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Volcanic history of the Imbrium basin: A close-up view from the lunar rover Yutu

Jinhai Zhang^a, Wei Yang^a, Sen Hu^a, Yangting Lin^{a,1}, Guangyou Fang^b, Chunlai Li^c, Wenxi Peng^d, Sanyuan Zhu^e, Zhiping He^f, Bin Zhou^b, Hongyu Lin^g, Jianfeng Yang^h, Enhai Liuⁱ, Yuchen Xu^a, Jianyu Wang^f, Zhenxing Yao^a, Yongliao Zou^c, Jun Yan^c, and Ziyuan Ouyang^j

^aKey Laboratory of Earth and Planetary Physics, Institute of Geology and Geophysics, Chinese Academy of Sciences, Beijing 100029, China; ^bInstitute of Electronics, Chinese Academy of Sciences, Beijing 100190, China; ^cNational Astronomical Observatories, Chinese Academy of Sciences, Beijing 100012, China; ^dInstitute of High Energy Physics, Chinese Academy of Sciences, Beijing 100049, China; ^eKey Laboratory of Mineralogy and Metallogeny, Guangzhou Institute of Geochemistry, Chinese Academy of Sciences, Guangzhou 510640, China; ^fKey Laboratory of Space Active Opto-Electronics Technology, Shanghai Institute of Technical Physics, Chinese Academy of Sciences, Shanghai 200083, China; ^gThe Fifth Laboratory, Beijing Institute of Space Mechanics & Electricity, Beijing 100076, China; ^hXi'an Institute of Optics and Precision Mechanics, Chinese Academy of Sciences, Xi'an 710119, China; ⁱInstitute of Optics and Electronics, Chinese Academy of Sciences, Chengdu 610209, China; and ^jInstitute of Geochemistry, Chinese Academy of Science, Guiyang 550002, China

Edited by Mark H. Thiemens, University of California, San Diego, La Jolla, CA, and approved March 24, 2015 (received for review February 13, 2015)

We report the surface exploration by the lunar rover Yutu that landed on the young lava flow in the northeastern part of the Mare Imbrium, which is the largest basin on the nearside of the Moon and is filled with several basalt units estimated to date from 3.5 to 2.0 Ga. The onboard lunar penetrating radar conducted a 114-m-long profile, which measured a thickness of ~5 m of the lunar regolith layer and detected three underlying basalt units at depths of 195, 215, and 345 m. The radar measurements suggest underestimation of the global lunar regolith thickness by other methods and reveal a vast volume of the last volcano eruption. The in situ spectral reflectance and elemental analysis of the lunar soil at the landing site suggest that the young basalt could be derived from an ilmenite-rich mantle reservoir and then assimilated by 10–20% of the last residual melt of the lunar magma ocean.

volcanic history | Imbrium basin | lunar rover Yutu |
lunar penetrating radar | Chang'e-3 mission

The surface of the Moon is covered by regolith, a mixed layer of fine-grained lunar soil and ejecta deposits, which is crucial to understanding the global composition of the Moon. The lunar regolith has also recorded the complex history of the surface processes, and it is the main reservoir of ³He and other solar wind gases. The thickness of the lunar regolith was estimated to be from 2 to 8 m in the maria and up to 8–16 m in the highland areas using various methods (1), including crater morphology (2, 3), seismology with low spatial resolution (4), radar wave scattering (5), and microwave brightness temperature (6). However, no in situ measurement of spectral reflectance, elemental compositions, lunar regolith thickness, or subsurface structures has been carried out.

The surface of the Moon is dominated with numerous large basins. They were formed about 3.9 Ga (7, 8), probably by the late heavy bombardment, and then filled with dark lava flows derived from partial melting of the lunar mantle, within a period mainly during 3.8–3.1 Ga (7). The Imbrium basin is the largest and was formed on Procellarum KREEP [potassium (K), rare earth elements (REE), and phosphorus (P)] Terrane (9), a unique terrain highly enriched in U, Th, and K radionuclides and other incompatible trace elements referred to as KREEP (10) and considered as the last residual melt of the Lunar Magma Ocean (11). The presence of the KREEPy materials, indicated by high concentrations of radionuclides U, Th, and K (9), around the rims of the Imbrium basin suggests that they are likely the basin-forming ejecta deposits. At least three main lava flows, dated from 3.5 Ga to 2.0–2.3 Ga (7, 12), have been recognized in Mare Imbrium with distinct FeO and TiO₂ concentrations (13, 14), which brought up interior information of this KREEP-rich terrain. The old and low-Ti basalt unit has been sampled by the Apollo 15 mission that landed at the eastern rim of the Imbrium basin. Information of other lava

flows in Mare Imbrium was obtained only by remote sensing from orbit. On December 14, 2013, Chang'e-3 successfully landed on the young and high-Ti lava flow in the northeastern Mare Imbrium, about 10 km south from the old low-Ti basalt unit (Fig. 1).

The lunar rover Yutu (named for the jade rabbit on the Moon in a Chinese fairy tale) was equipped with an active particle-induced X-ray spectrometer (APXS), a visible to near-infrared (450–945 nm) imaging spectrometer and short-wave infrared (900–2,395 nm) spectrometer (VNIS), and a lunar penetrating radar (LPR), accompanied by a stereo camera and a navigating camera. Originally, the mission planned to have the lunar rover measure chemical and mineral compositions of the lunar soil and various types of ejecta rocks and to carry out a LPR profile of the lunar regolith and subsurface structures in the first 3 mo. The mission was scheduled to extend up to 1 y and to explore the old low-Ti lava flow ~10 km north. Unfortunately, some of Yutu's mechanical parts failed to move just before the rover prepared for sleeping at the end of the second month due to unknown faults probably in the control system. During the first 2 mo, Yutu successfully carried out two APXS and four VNIS analyses of the lunar soil and performed a 114-m-long LPR profile along the rover track in the landing area (Fig. 2). These in situ measurements

Significance

After the Apollo and Luna missions, which were flown about 40 years ago, the Moon was explored only from orbit. In addition, no samples were returned from the young and high-FeO and TiO₂ mare basalt in the northern Imbrium basin. Such samples are important to understand the formation and evolution of the Procellarum KREEP [potassium (K), rare earth elements (REE), and phosphorus (P)] terrain, a key terrain highly enriched in radioactive nuclides. The Chang'e-3 mission carried out the first in situ analyses of chemical and mineral compositions of the lunar soil and ground-based measurements of the lunar regolith and the underlying basalt units at this specific site. The lunar regolith layer recorded the surface processes of the Moon, whereas the basalt units recorded the volcanic eruption history.

Author contributions: Y.L., C.L., Y.Z., J. Yan, and Z.O. designed research; J.Z., W.Y., S.H., Y.L., and W.P. performed research; G.F., W.P., Z.H., H.L., J. Yang, E.L., and J.W. contributed new reagents/analytic tools; J.Z., W.Y., S.H., Y.L., S.Z., B.Z., Y.X., and Z.Y. analyzed data; and J.Z., W.Y., S.H., Y.L., and Y.X. wrote the paper.

The authors declare no conflict of interest.

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¹To whom correspondence should be addressed. Email: linyt@mail.igcas.ac.cn.

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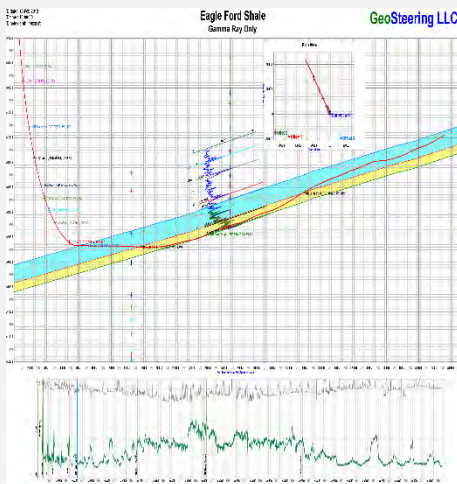
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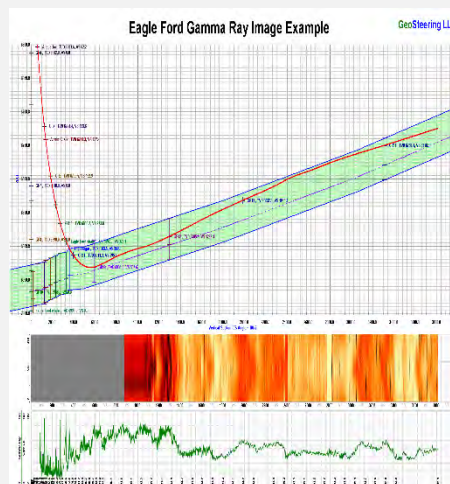
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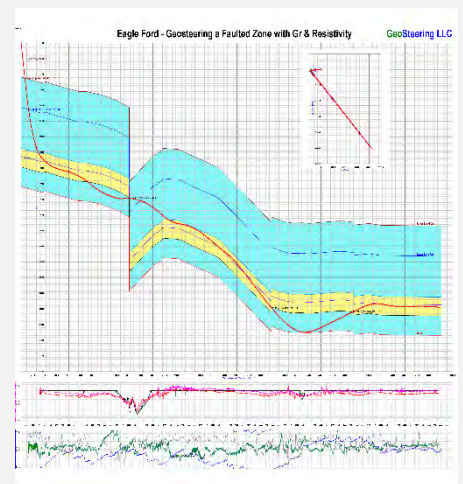
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estimated from the corrections between the compositions and the reflectance spectrum parameters of the Apollo soils (*SI Appendix*, Fig. S9), which are 18.9 wt % (18.7–19.5 wt %) and 6.6 wt % (5.3–9.0 wt %), respectively. The estimated TiO₂ contents are somewhat higher than the APXS analyses, likely due to the shadow of the rough lunar soil surface (*SI Appendix*, Fig. S8). The estimated FeO contents are consistent with the APXS analyses within analytical uncertainties. The space weathering effect (darkening and reddening of the spectra) of the lunar soil can be indicated by the optical maturity index (18), which ranges from 0.098 to 0.158 with an exception of 0.312 for CD005 (*SI Appendix*, Table S2). The higher values of the optical maturity index and the spectral reflectance of CD005 may be due to blowing the top dusts off the lunar surface by the rocket during the descent of Chang'e-3 because CD005 is located closest to the lander (*SI Appendix*, Table S2). The optical maturity index values of CD006–CD008 suggest submature or mature lunar soil, undistinguishable from the Apollo mare soils. This is inconsistent with the rocky surface and the presence of abundant rocks in the lunar soil at the landing site (*SI Appendix*, Fig. S2), which is suggestive of a young age for the Chang'e-3 landing site. However, it has been noticed that there was no clear correlation between the optical maturity index and the ages of the Apollo mare soils (18).

Lunar Regolith Layer Thickness

The lunar regolith layer and the underlying basalt units at the landing site have been detected by the onboard LPR, which has two frequency channels, i.e., channel 1 at 60 MHz and channel 2 at 500 MHz. The echoes were processed using preprocessing, migration and time–frequency analysis methods that are commonly used in seismic exploration (*SI Appendix*). Fig. 4 displays the LPR profile of channel 2 with a total distance of 73 m along the Yutu's track. The relative dielectric constant ϵ_r increases from 2.1 at the surface to 6.5 at a depth of 10 m, based on the function of lunar soil density with depth (19) and a model of the lunar regolith. We propose that the homogeneous uppermost sublayer contains 5.7 vol % of basaltic blocks determined from the high-resolution images of the landing camera (*SI Appendix*, Fig. S2). Then, we assume that the rock debris content of the lunar regolith increases with depth to 100% at 10 m deep. The high-frequency LPR echoes clearly exhibit an uneven bottom of the lunar regolith, and the depth varies from 2.2 m to 5.4 m with a median value of ~5 m (Fig. 4).

The thickness of the lunar regolith layer can also be constrained by high-resolution morphology of craters taken by the landing camera of Yutu. There are two large craters in the landing region (Fig. 2). The larger one (crater A, 18 m in diameter) is blocky, indicative of penetrating the regolith layer and excavating the underlying bedrock, whereas the other (crater B, 13.7 m in diameter) has few blocks, and hence, it is inside the regolith layer. The depths of both craters were calculated to 3.7 m and 4.9 m using the impacting model of the lunar regolith (www.lpi.usra.edu/lunar/tools/lunarcatercalc/). The depths of 3.7 m and 4.9 m can be referred to as the range of the regolith layer thickness, which is consistent with the LPR measurement. The ground-based LPR measurement of the lunar regolith layer thickness clarifies the wide range in previous estimations from 2.6 m (20) to 7.5 m (21) using the crater morphology method, or 6.2 ± 2.0 m (5) and 6.4 ± 1.7 m (22) based on lunar radar sounder data, in the same young, high-Ti region of Mare Imbrium. A thickness of 8.5–12.2 m of the regolith layer in the Apollo highlands was measured with seismology (4), but the lunar regolith layer thickness of 2–8 m at the Apollo mare sites was estimated (1). In addition, a much thicker regolith (8–32 m, with a median depth of 11 m) was estimated in three Eratosthenian mare areas (2.0–2.5 Ga) (23). Based on the positive correlation between regolith layer thickness and age of the

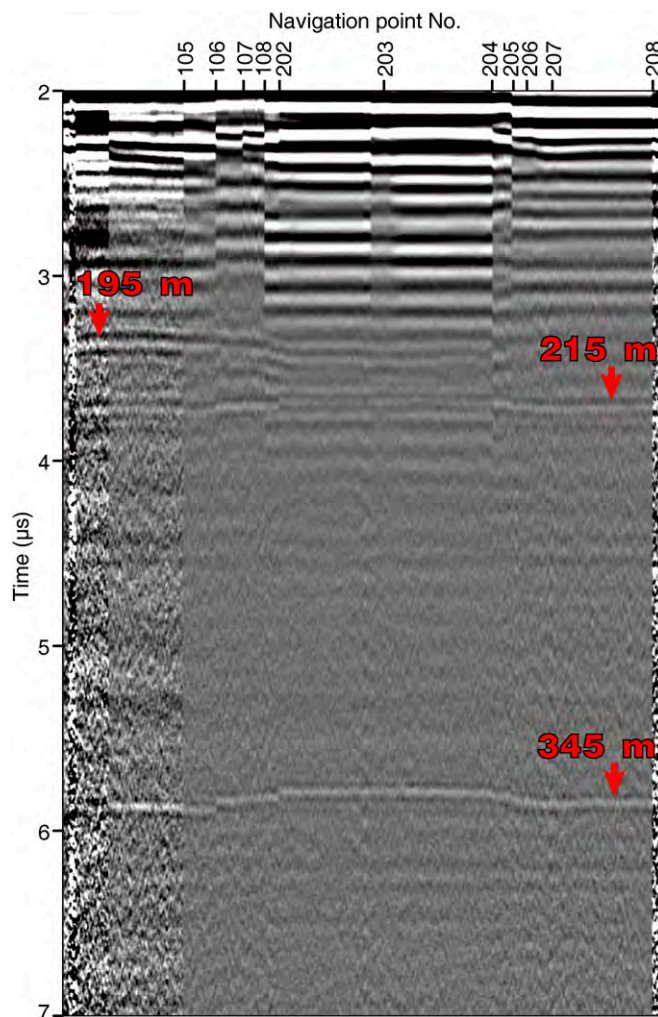


Fig. 5. The LPR profile of channel 1.

surface (5), the regolith layer thickness at the young Chang'e-3 landing site could be considered to be the minimum of the global surface because of its young age of 3.0 Ga (7) or 2.0–2.3 Ga (12). The thick regolith layer at the Chang'e-3 landing site may be partially attributed to the ejecta deposits from a crater ~450 m in diameter, which locates ~60 m west to the lander. However, the high-resolution descending images reveal no recognizable modification of the distribution pattern of small craters on the surface (Fig. 2). The thickness of ejecta deposits must be less than 20 cm; otherwise, the craters <0.5 m in diameter would have been buried. Hence, the echoes detected by the LPR are mainly from the regolith layer rather than the ejecta deposits of the nearby crater. The mean thicknesses of the lunar regolith layer at Apollo mare sites were generally reported to 2–8 m (1), and the lower range could be somewhat underestimated because these sites are significantly older (3.8–3.1 Ga). This is also confirmed by the blocky nature of the Chang'e landing site, which suggests a thinner regolith layer than the Apollo sites.

In Fig. 4, the uppermost sublayer (~0.7 m in thickness) can be observed. The presence of this sublayer is independent of the model of the lunar regolith we used (*SI Appendix*, Figs. S16 and S17). This homogeneous sublayer likely represents the well-plowed-up top zone of the lunar regolith caused by small impacts. The measured thickness of ~0.7 m of the sublayer is consistent with that of the homogeneous zone (~60 cm) at the top of Apollo 17 deep drill cores (1).



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Underlying Basalt Units

The underlying basalt units were probed by the low-frequency LPR echoes. Three obvious reflectors can be observed at depths of 195 m, 215 m, and 345 m (Fig. 5), using $\epsilon_r = 6.5$ (1). These reflectors likely indicate buried regolith layers formed during flooding hiatus of lava flows in the northeastern Mare Imbrium. This is consistent with the presence of three main types of basalt units in Mare Imbrium, based on variation of FeO and TiO₂ contents (13, 14). These basalt units were dated 3.3–3.5 Ga, 3.0 Ga, and 2–2.5 Ga (7, 12, 24). The lava flows originated from the vent area near the crater Euler (over 700 km south to the landing site) and extended north 1,200 km, 600 km, and 400 km from older low-Ti basalts to the overlying high-Ti flows (*SI Appendix, Fig. S1*). The buried regolith layers between these basalt units are expected, produced by asteroid impacts during the time periods between the eruptions of the basalt units, which are likely responsible for these significant reflection horizons. Hence, the thickness of these three basalt units is 195 m, 20 m, and 130 m, from the top to the bottom, respectively.

Based on the effects of lava flow on lunar crater populations, the thickness of lava flows in the southwest Mare Imbrium were determined to 30–60 m (25) or six flow units with an average of 32–50 m (26). The much thicker uppermost basalt unit determined by Yutu's LPR suggests a sequence of lava flooding events (e.g., refs. 3–6) during a short period, and the individual flows cannot be detected by the LPR because of little regolith produced in the short hiatus. The total thickness of the top two basalt units (both high Ti) can be constrained by morphology of small craters with diameters of 0.4–4.0 km. The distribution of the TiO₂ contents of the rims and proximal ejecta deposits of these craters determined from compositional mapping data (14) revealed whether or not the craters have penetrated the overlying layer and excavated the bedrock beneath. The smallest craters with a diameter of 0.4–1.0 km show a peak at ~5.0 wt % TiO₂ with a low-TiO₂ tail, whereas that of large craters has a low-TiO₂ peak at ~2.5 wt % (*SI Appendix, Fig. S7*). The high-TiO₂ peak is similar to the in situ APXS analyses and the remote sensing results (27, 28). The low-TiO₂ tail of the smallest craters is due to excavating the underlying low-Ti basalt unit, as indicated by the low-TiO₂ peak for the larger craters (1–4 km in diameter) and the presence of old low-Ti basalt about 10 km north to the landing site. This result sets up a range of the total thickness of the upper two basalt units, from significantly larger than 109 m (depth of a crater with 0.4-km diameter) to obviously less than 270 m (depth of a crater with a 1-km diameter). This is consistent with the total thickness (210 m) of the upper two reflectors of LPR. Two subsurface layers in the eastern Mare Imbrium (36.0° N, 15.3° W) have been detected by lunar radar sounder on the Kaguya spacecraft, with depths of 250 m and 454–460 m (29). The footprint of the lunar radar sounder is located on the old low-Ti basalt unit (29), and the first reflector could be assigned to the third echo of Yutu's LPR. However, the depth of 250 m is nearly twice of the thickness of the old low-Ti basalt unit determined by Yutu's LPR. In addition, the deeper layer at 454–460 m measured by Kaguya spacecraft has not been detected by Yutu's LPR.

Conclusions

The chemical compositions of the lunar soil at the Chang'e-3 landing site were determined in situ by APXS, including trace elements of Zr, Sr, Y, and Nb. Its higher FeO (21.3–22.1 wt %) and TiO₂ contents (4.0–4.3 wt %) and lower Al₂O₃ content (10.5–11.5 wt %) compared with the Apollo and Luna soil

samples suggest a new type of basalt beneath, which has not yet been sampled in previous missions. The nearly flat KREEP-normalized abundance pattern of K, Zr, Y, and Nb suggests assimilation of the underlying KREEP material. The mineral abundances decoded from the ground-based measurements of the reflectance spectra of the lunar soil are similar to Apollo mare soil.

The thickness of the lunar regolith layer was measured with the LPR to be ~5 m. This is significantly larger than what we expected for a young mare site dated about 2.0–3.0 Ga (7, 12). Such a thick lunar regolith layer is unlikely attributed to ejecta deposits from a nearby crater because no significant modification on the surface was observed in the high-resolution images taken by the descending camera. In addition, an upper sublayer of ~70 cm was detected, which has a feature of lacking rocks. This argues against a detectable layer of the ejecta deposit along the lunar rover's track. Our LPR imaging results show that the thickness of the global lunar regolith layer was underestimated in previous studies (e.g., refs. 1, 20).

The presence of three basalt units detected by the LPR is consistent with the distribution of three main basalt units in the northern Mare Imbrium (*SI Appendix, Fig. S1*). The thickness of 195 m of the top basalt unit indicates no significant cessation of volcanic eruption in the Imbrium basin until 2.0–3.0 Ga, different from a gradual cessation of magma eruption in the Imbrium basin (30). This could be at least partially attributed to high concentrations of radionuclides in the Procellarum KREEP Terrane. The high-Ti basalt was probably formed via partial melting of ilmenite-rich lunar mantle reservoir, followed by ascending to the Imbrium basin with 10–20% contamination of the KREEP material.

Methods

The chemical compositions of the lunar soils were determined from the APXS analyses, calibrated with two analyses of the onboard basaltic working reference and a set of standards that were measured in laboratory. All of the original APXS data are in 2B level and have been corrected for background and peak overlapping (for data processing details, see *SI Appendix, S2*).

The mineral abundances, optical maturity index, and the contents of FeO and TiO₂ of the lunar soil were achieved from the VNIS data in 2B level. The VNIS spectra were combined from those measured separately by the VIS/NIR imaging spectrometer (450–945 nm) and the shortwave IR (SWIR) (900–2395 nm), and they were reduced to repair bad lines and make flat field corrections and converted to reflectance. The Modified Gaussian Model was applied to decode the VNIS spectra, and the Apollo mare soils data were used to calibrate the results (for data handling, see *SI Appendix, S4*).

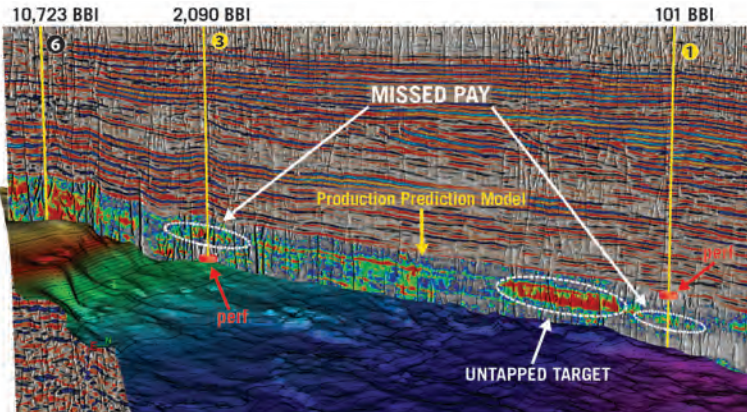
The lunar regolith layer was detected by the high-frequency channel (500 MHz, channel 2), and the depths of underlying basalt units were detected by the low-frequency channel (60 MHz, channel 1) of LPR on the Yutu rover. The 2B level data of channel 2 were processed using spherical spreading amplitude compensation and filtered by the second-order Butterworth filter. The time–frequency analysis was applied to the channel 2 data. The depths of the lunar soil sublayer and the lunar regolith bottom were estimated using the migration method. For channel 1, the 2B level data were filtered by band-pass Butterworth filter and 2D median filter (for full technical details, see *SI Appendix, S5*).

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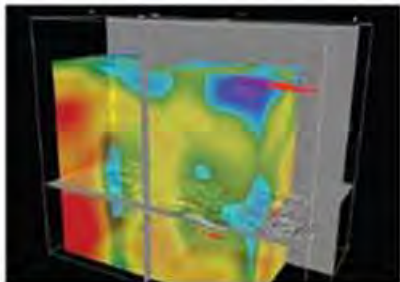
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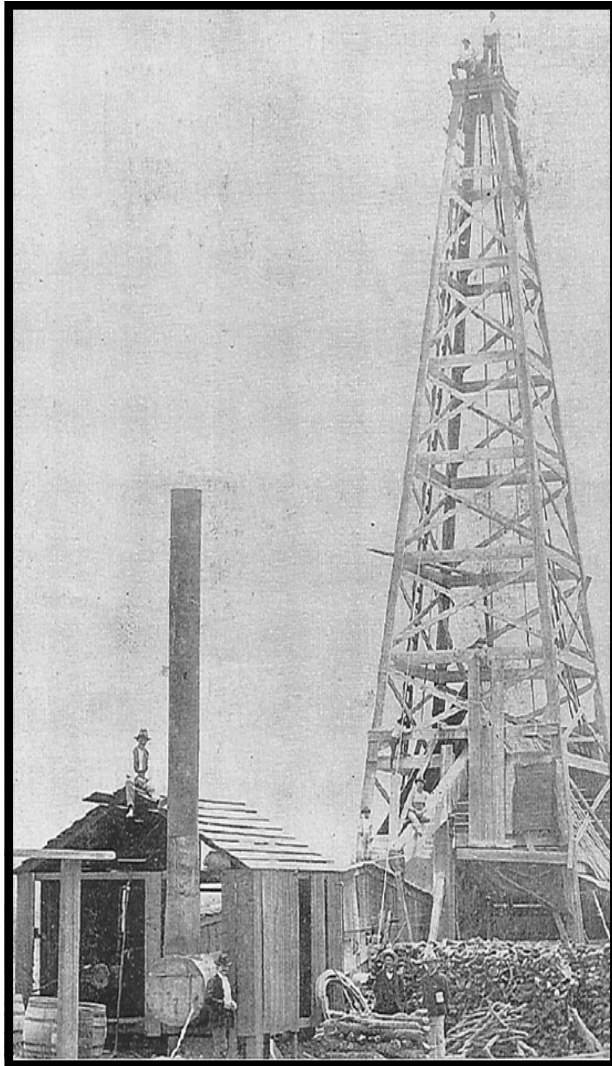
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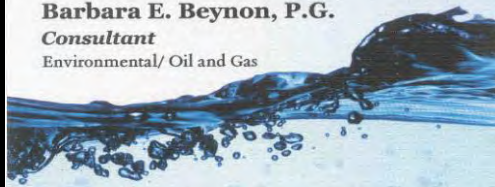
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Geologist
600 Leopard St, Ste 706
Corpus Christi, TX 78401

dkbecker1137@sbcglobal.net

361-884-3613

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(361) 774-3369 Bbeynon77@yahoo.com



**Dawn S. Bissell
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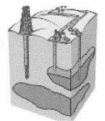
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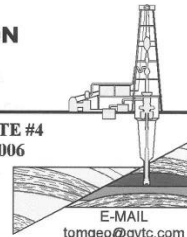
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Richard M. Parker
Consulting Geologist

Adjunct Professor - Geosciences
Texas A&M University - Kingsville
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Doug Sartoris

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TX LICENSE #2289

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1330 Post Oak Boulevard
Suite 600
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WILLIAM A. WALKER, JR.
Certified Petroleum Geologist
bwalker@stalkerenergy.com

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SEBASTIAN P. WIEDMANN
GEOLOGIST

WILSON PLAZA WEST
606 N. CARANCAHUA, SUITE 500
CORPUS CHRISTI, TEXAS 78401

Mobile (361) 946-1430
swiedmann_geo@mail.com



Dave Willis
Onshore Sales

Main +1 713 789 7250
Direct +1 281 781 1035
Mobile +1 281 543 6189
Fax +1 713 789 7201
dave.willis@iongeo.com

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ROLF WOODS
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